# Driving of Strong Nightside Reconnection and Geomagnetic Activity by Polar Cap Flows: Application to CME Shocks and Possibly Other Situations

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## 29 Abstract

30 Previous studies have shown that dynamic pressure impacts (e.g., shocks 31 initiating CME storms) with southward IMF promptly lead to strong auroral nightside 32 activity and concurrent poleward expansion (indicating strong nightside reconnection), 33 and strong enhancements in convection and currents. Here we use a combination of 34 ground-based ASI and radar observations to further describe this response, to address 35 what is driving the strong activity, and to suggest similar driving in other situations. 36 Consistent with some previous studies, we find that shock driven auroral activity and 37 poleward expansion resembles a substorm, but starts from an already broad MLT sector 38 without much subsequent azimuthal expansion and without classical brightening of the 39 equatorward-most arc. We furthermore find a large enhancement of meso-scale 40 ionospheric polar cap flows heading towards the nightside separatrix immediately after 41 shock impact. Recent studies have shown that such enhanced flows often cross the 42 separatrix leading to plasma sheet flow bursts, poleward boundary intensifications (PBIs), 43 streamers, and poleward motion of the polar cap boundary from reconnection. Thus these 44 flow enhancements, which must extent outward along field lines from the ionosphere, are 45 an attractive candidate as the driver for the almost immediate strong auroral, current, and 46 reconnection activity resulting from shock impact. We also discuss and present some 47 evidence that this phenomenon may be more general, leading to similar oval responses 48 without a shock impact, including during and following the expansion phase some 49 substorms. These suggestions could lead to some possibly fundamental questions, such 50 as when do polar cap convection enhancements lead to a substorm growth phase versus 51 leading directly to strong polar expansion of, and strong activity along, the auroral oval 52 field line region?

53 Keywords: Magnetic storms; aurora; convection; polar cap; substorm; reconnection

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### 55 **1. Introduction**

56 Abrupt enhancements of solar wind dynamic pressure (P<sub>dyn</sub>), such as the shocks 57 that initiate coronal mass ejection (CME) storms, cause dramatic effects when they occur 58 under southward interplanetary magnetic field (IMF) conditions. They drive nearly 59 immediately strong auroral activity, poleward expansion of the auroral oval that can reach as much as 10<sup>o</sup> in latitude, and strong enhancement in global convection, 60 61 ionospheric currents, and Region 1 and Region 2 field-aligned currents (Boudouridis, 62 2003; Lyons et al., 2016; Zesta et al., 2000). The rapid poleward expansion of the 63 nightside auroral oval implies strong nightside reconnection in the presence of a nearly 64 simultaneous increase in the strength of convection.

65 The nightside auroral activity resembles a substorm, but initiates over a 66 substantially broader range of MLT without much subsequent azimuthal expansion as 67 poleward expansion occurs over a broad longitudinal range (Chua et al., 2001; Lyons et 68 al., 2000; Zesta et al., 2000). These and other features, such as the lack of brightening of 69 an equatorward arc, have lead to the suggestion that the disturbance is different from 70 substorms (Liou et al., 2003; Yue et al., 2013). However, other studies have indicated 71 that the disturbance may be a substorm (Kokubun et al., 1977; Lyons, 2005; Zhou and 72 Tsurutani, 2001), and dipolarizations as occur during substorms are seen at 73 geosynchronous orbit in response to solar wind dynamic pressure impacts on the 74 magnetosphere (Lee et al., 2005; Lee and Lyons, 2004).

In this paper, we describe the dramatic nightside response using modern groundbased all-sky imager (ASI), radar, and low-altitude spacecraft observational capabilities, and we address what is driving the strong auroral, current, and reconnection activity. We take particular advantage of the 17 March 2013 storm, a CME-driven event initiated by a shock that impacted the magnetosphere at 06 UT (Baker et al., 2014). There has been much interest in ring current particle injections (e.g., *Gkioulidou et al., 2014*) and radiation belt electrons (Hudson et al., 2015; Li et al., 2014, 2015) for this event. Of 82 importance for the current study is the excellent radar and auroral observation coverage at 83 times just before and just after the shock impact (Lyons et al., 2016; hereafter referred to 84 as Paper 1), allowing for excellent evaluation of the effects of the shock. Consistent with 85 some previous studies, we find that the onset of the shock-driven activity appears to be 86 very different from the onset of a substorm. However, we find that the activity following 87 shock impact appears to be driven by enhancements of meso-scale flows along polar-cap 88 field lines, and this has important similarities to what has been suggested to drive 89 prolonged activity during the expansion phase of some substorms. This, and additional 90 evidence we present here, suggest that the driving of activity we show here may apply 91 more generally than just to the impact of P<sub>dyn</sub> enhancements.

### 92 **2. Observations**

93 Observations for the magnetic storm on 17 March 2013 in Paper 1 show that the 94 shock impact with concurrent southward IMF immediately drove dramatic poleward 95 expansion of the poleward boundary of the auroral oval (implying strong nightside 96 reconnection), strong auroral activity, and strong penetrating mid-latitude convection and 97 ionospheric and field-aligned currents. Figure 1 shows, from top to bottom, the WIND 98 solar wind dynamic pressure Pdyn, the OMNI interplanetary magnetic field (IMF) as 99 propagated to the dayside magnetopause, the SuperMAG (Gjerloev, 2012) ground 100 magnetometer upper U and lower L auroral magnetic index, and the SuperMAG ring 101 current index for all MLT and within the dusk, noon, dawn, and midnight sectors. The 102 SuperMAG indices are based on the traditional AU, AL, and SYMH indices, but with 103 many more stations. The OMNI IMF data is shown further shifted by  $\sim 10$  minutes so that 104 the shock impact time agrees with the 0600 UT impact time seen by the dayside ground 105 magnetometers. The WIND P<sub>dvn</sub> is shown because of data gaps in the OMNI P<sub>dvn</sub> around 106 the time of the shock and is further shifted by ~30 min relative to the IMF data to agree 107 with the time of shock impact.

108 The U and L indices reflect immediate and large increases in ionospheric current 109 that were driven when the shock impacted the magnetosphere, as indicated by the large increase in Pdyn. The lower panels show the strong increases in Region 1 and Region 2 110 111 currents that were seen from AMPERE magnetic perturbations observed along Iridium 112 satellite trajectories during the 10 min interval immediately before, and the second 10 113 min interval after, the shock impact. Red and blue shadings give upward and downward 114 FACs, respectively, obtained from the curl of fits to the magnetic perturbations (Waters 115 et al., 2001).

116 Figure 2 shows mergers of auroral images over Canada from the array of 117 THEMIS ASIs (Mende et al., 2008) overlaid with line-of-sight (LOS) flow velocities 118 from the Super Dual Auroral Radar Network (SuperDARN) radars from the period a few 119 min before to 18 min after the P<sub>dvn</sub> impact. The ASI images are displayed so as to 120 emphasize the radar echoes within the polar cap from the pair of PolarDARN radars 121 (Koustov et al., 2009) at Inuvik to the west and Rankin Inlet to the east. Observations 122 from the lower latitude SuperDARN radars have been emphasized in Paper 1. The top 123 three panels of Figure 2 show a narrow band of moderately active aurora that lay along 124 the poleward boundary of the evening-to-midnight auroral oval prior to the shock impact. 125 This activity, and the lower than average location of the auroral poleward boundary (magnetic latitude  $\Lambda \sim 68^{\circ}$  -70°), likely resulted from the pre-storm southward IMF of a 126 127 few nT. The dramatic enhancement in activity and poleward expansion of the auroral 128 oval that initiated almost immediately after the shock impact can be seen over several 129 hours of MLT in the next four panels (0602 to 0611 UT) of Figure 2, as well as in movie 130 S1 in paper 1. This includes a rapid, large poleward expansion of the auroral oval over a 131 broad MLT range.

The enhancement and poleward expansion of the aurora also shows vividly in the
auroral observations from the Special Sensor Ultraviolet Spectrographic Imager (SSUSI)
(http://ssusi.jhuapl.edu/) onboard polar-orbiting Defense Meteorological Satellite

Program (DMSP) spacecraft in Figure 3. SSUSI measures FUV emissions in five spectral
bands simultaneously (Paxton et al., 2017). Algorithms have been applied to properly
remove the dayglow, and the LBH emission intensities shown in Figure 3 reflect
precipitating auroral electron energy flux (Paxton et al., 1992, 1998, 1999).

139 The auroral images in Figure 3 are essentially keograms made along the trajectory 140 of a spacecraft from horizon to horizon imager scans in the direction normal to the 141 trajectory. The DMSP spacecraft had extremely fortuitous crossings of the nightside 142 auroral oval in the southern hemisphere, F17 crossing just 12 to 4 min before the 0600 143 UT shock impact, and F18 and F16 crossings from 1 to 14 min after the impact. Strong 144 auroral enhancement and poleward expansion can be seen as the F18 and F16 spacecraft 145 passed the dawnside oval at  $\sim 0602$  to 0604, just 2 to 4 min after shock impact, and the F16 observations show that the oval expanded poleward by ~5° in latitude by ~0606-146 147 0607 UT. The poleward expansion can also be seen in the precipitating electron energy 148 fluxes measured by DMSP and plotted along the spacecraft trajectory (the full 149 precipitating particle energy spectra from the F16 and F17 crossings are shown in Figure 150 3 of Paper 1).

151 As seen from the SuperDARN observations in Figure 2, LOS velocities were 152 moderately strong within the nightside polar cap during the period preceding the shock 153 impact. They had an anti-sunward LOS component, and show considerable meso-scale 154 structure with longitudinal scales of ~0.5 to 1 hr in MLT. Immediately following the 155 impact, the anti-sunward LOS ionospheric flows seen by the radars and their meso-scale 156 structure enhanced considerably within the polar cap region encircled by the ellipses, and 157 peak values above 600 m/s can be seen directed towards the nightside auroral oval. These 158 flows can be seen down to magnetic latitudes  $\Lambda$  of ~ 77-78°, just a few degrees poleward 159 of the expanded auroral oval. (The few radar echoes at  $\Lambda \sim 73-76^{\circ}$  are echoes for the E-160 region, where speeds are substantially reduced from the electric field drift speed). The 161 polar cap flows and auroral activity decreased approximately simultaneously starting at 162 0614 UT, as the IMF briefly turned northward. A similar enhancement of polar cap flows 163 and their meso-scale structure can be seen in SuperDARN observations in Figure 3 from 164 the southern hemisphere SuperDARN radars. In Figure 3, enhanced polar cap flows seen 165 by SuperDARN near midnight and near 04 MLT appear to impact the poleward boundary 166 of the auroral oval. The inference is supported by the flows measured by the DMSP F17 167 and F16 spacecraft in the direction normal to the spacecraft trajectory, these flows being 168 shown by the violet bars in Figure 3.

169 Figure 4 shows the poleward expansion of the oval in the southern hemisphere 170 based on the precipitating electron energy fluxes from both the Polar Operational 171 Environmental Satellites (POES) and the DMSP spacecraft. Though the two post-shock 172 POES passes crossed the nightside auroral oval several minutes after the shock-related 173 activity started to subside, they clearly show that the oval poleward boundary had 174 expanded to  $\Lambda \sim 78^{\circ}$ , which is  $\sim 10^{\circ}$  poleward of where it was prior to the shock impact. 175 Figure 4 also includes the polar cap flows seen by SuperDARN at times during the period 176 of rapid auroral poleward expansion. The enhanced flows can clearly be seen to have 177 extended from the polar cap into the region that became occupied by the expanded oval. 178 As in Figure 3, the flows from the polar cap appear to be directed so as to impact the 179 auroral poleward boundary, though the lack of direct observation of flows impacting the 180 auroral poleward boundary prevents definitive determination of whether this occurred.

181 Before addressing the possible role of the enhanced polar cap flows in the 182 magnetosphere responses to dynamic pressure impacts and their possible generalization 183 to other activity, we discuss observations from two additional storm events initiated by 184 impacts of CME shock increases in P<sub>dvn</sub> under southward IMF at times appropriate for 185 ASI viewing over North America. The first storm was a CME storm on 19 February 186 2014 (Durgonics et al., 2017). Figure 5 shows the OMNI interplanetary magnetic field 187 (IMF) and P<sub>dyn</sub> as propagated to the dayside magnetopause, the SuperMAG ground 188 magnetometer upper U and lower L auroral magnetic index, and the SuperMAG ring 189 current index for all MLT and within the dusk, noon, dawn, and midnight sectors for this 190 event. The OMNI data have been shifted by ~12 minutes so that the shock impact time 191 agrees with the 0348 UT impact time seen in the ring current indices. As for the 17 192 March 2013 storm, the U and L indices show immediate and large increases in 193 ionospheric current that where driven as the shock impacted the magnetosphere, as 194 indicated by the large increase in P<sub>dyn</sub>

195 Figure 6 shows mergers of auroral images over Canada from the THEMIS ASIs 196 overlaid with line-of-sight (LOS) flow velocities from the SuperDARN radars from the 197 period a few min before to 30 min after the shock impact. Moderately active aurora can 198 be seen prior to the impact in the first two panels, the activity likely associated with the 199 substantially negative IMF B<sub>z</sub> during that period. Poleward expansion of the oval by a 200 few degrees in latitude and an enhancement in activity can be seen over several hours of 201 MLT during the 20 min period after the shock impact in the next five panels (0350 to 202 0410 UT) of Figure 2. It can be seen that the rapid, large poleward expansion of the 203 auroral oval occurred over a broad MLT range. The SuperDARN observations in Figure 204 6 (encircled by the yellow ellipse) show an enhancement of LOS velocities directed anti-205 sunward within the nightside polar cap during the period of auroral poleward expansion. 206 As indicated by yellow arrows, the observations at 0350 and 0352 UT show enhanced 207 LOS flows at 20-22 MLT (the midnight meridian is given by the blue line) at A's from 208  $\sim 68^{\circ}$  to 72°. These flows appear to be enhanced over those prior to shock impact, to 209 impinge upon the nightside auroral poleward boundary, and to extend to within the 210 auroral oval. We note that the numbers of radar echoes in this region prior to the shock 211 impact were limited, preventing definitive determination of whether the flows were 212 enhanced by the shock impact.

Enhancement of anti-sunward polar cap flows also occurred in the southern hemisphere for this storm. This can be seen from the SuperDARN LOS flows in Figure 7, the region of the polar cap echoes being quite near the magnetic pole. The poleward expansion of the auroral oval is seen quite clearly in Figure 7 to have been  $\sim 5^{\circ}$  in latitude by 0357 UT (9 min after shock impact), based on the precipitating electron energy fluxes from the POES spacecraft that are shown in the lower-right panel.

219 The third event with southward IMF had an onset at 0826 UT at 5 April 2010 and 220 was studied by Yue et al. (2013) (See their Figure 7-9) and Loto'aniu et al. (2015). (This 221 was the storm that brought down the Galaxy 15 satellite for about 1 year.) Quick and 222 very strong increases in nightside, anti-sunward polar cap flows with meso-scale structure 223 were seen also for this event, as shown by the nightside SuperDARN LOS velocities in 224 Figure 8. For this event, there were THEMIS spacecraft nicely positioned near the northern hemisphere outer boundary of the plasma sheet near midnight at  $X_{GSM} = -11.3 R_{E}$ . 225 226 These spacecraft saw the initiation of plasma sheet flow bursts just ~2-3 min after the 227 dynamic pressure impact, consistent with the entry of the enhanced meso-scale flow from 228 open polar cap field lines to the plasma sheet. (Note that ionospheric flows impacting the 229 poleward boundary of the auroral oval would be expected to extend outward along field 230 lines into the magnetosphere and impact the outer boundary of the plasma sheet at all X 231 distance along the open-closed field line boundary that lie earthward of the distant 232 magnetic X-line. This is assuming that magnetic field lines are approximately 233 equipotentials and that the poleward boundary of the auroral oval and outer boundary of 234 the plasma sheet both lie along the boundary between open and closed field lines.)

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## 3. Possible role of enhanced polar cap flows

Enhancements in polar cap convection are often viewed as leading to substorm growth phase conditions (Juusola et al., 2011; McPherron, 1970). However, the observations here indicate that, under some conditions, enhancements in polar cap convection can lead to dramatic poleward expansion of, and activity within, the auroral oval, a response that is more like a substorm expansion phase than a substorm growth phase. Based on these observations, we suggest that the enhanced meso-scale polar cap flows heading towards the poleward boundary of the nightside auroral oval may be important for the enhancement of auroral activity and auroral poleward expansion associated with shock impacts. We also address evidence that this phenomenon may be more general than being only a response to abrupt  $P_{dyn}$  increases.

246 A large number of studies have shown that such flows often can cross the auroral 247 poleward boundary, and lead to plasma sheet flow bursts, intensifications along the 248 auroral poleward boundary (PBIs), and auroral streamers (de la Beaujardière et al., 1994; 249 Lyons et al., 2011; Nishimura et al., 2010; Ohtani and Yoshikawa, 2016; Pitkänen et al., 250 2013; Shi et al., 2012; Zou et al., 2014). This corresponds to externally driven localized 251 reconnection, assuming the auroral poleward boundary is approximately collocated with 252 the open-closed magnetic and flows cross that boundary. Driving of PBIs and streamers 253 by the enhanced polar cap flows may explain the enhancement in auroral activity 254 following the shock impact, which has been found in Paper 1 to consist of streamers. 255 Furthermore, poleward expansion of the auroral poleward boundary has been seen in 256 PBIs that result from incoming meso-scale flows from the polar cap (Zou et al., 2014, 257 2015). This indicates that poleward expansion of the auroral oval can be a result of the 258 reconnection and PBIs triggered by incoming polar cap flows.

259 To account for the poleward expansion of the oval that extends over several hours 260 of MLT and brings the auroral poleward boundary poleward by up to several degrees in 261 latitude, enhanced meso-scale flows would have to impact the auroral poleward boundary 262 over the entire longitudinal extend of the auroral poleward expansion and throughout the 263 period of that expansion. Consistent with this, as can be seen from Figures 2, 6, and 8, 264 enhanced flows with meso-scale structuring are seen over ~7-8 hours of MLT within the 265 nightside polar cap and heading towards the auroral poleward boundary. Furthermore, in 266 our cases, the flow initiates as the poleward boundary starts to move poleward. The flows 267 continue as the boundary continues to move poleward, and the flows reduce in magnitude 268 as the poleward motion of the oval ceases and activity decreases within the oval.

269 A limitation in our observations is that we do not have radar echoes just poleward 270 of the auroral poleward boundary at the time of the shock impact for the 17 March 2013 271 or 5 April 2010 storms, though we do have a limited number of such echoes for the 19 272 February 2014 storm. The activity along the auroral poleward boundary just prior to 273 shock impact indicates that enhanced meso-scale flows were impinging on the polar cap 274 boundary prior to the shock impacts. However, it is necessary for such flows to increase 275 with the shock impact as we have seen for the flows a few degrees further poleward in 276 order to account for the nearly simultaneous flow and auroral activity enhancement. 277 While do have the limited direct observation of the enhancement of these flows for the 19 278 February 2014 storm, as well as direct evidence from the increases observed deeper in the 279 polar cap, we cannot say for certain that, in general, flows adjacent to the polar cap 280 boundary increase at onset. Furthermore, the enhancement in field-aligned currents by 281 the shock compression would be expected to contribute to the shock-associated auroral 282 brightening, though we are not aware of evidence that such currents could contribute to 283 the poleward expansion of the aurora oval following shock impact.

284 In this paper, we have focused on P<sub>dyn</sub> increases that initiate CME-driven 285 geomagnetic storms. However, substantially enhanced meso-scale flows occur under 286 other conditions as well, and they may also drive enhanced aurora activity and poleward 287 expansion during these conditions. For example, in one of our earlier substorm studies 288 (Lyons et al., 2011), we noted unexpected observations from a limited number (8) of 289 substorm onsets that had good radar coverage within the polar cap region poleward of 290 substorm expansion phase activity. We found that four of the onsets were followed by 291 prolonged periods of strong ionospheric flow channels directed toward the polar cap 292 boundary from the polar cap. The expansion phases after these onsets had large auroral 293 poleward expansion and prolonged periods of PBI and streamer activity 294 contemporaneous with the enhanced equatorward-directed flow channels. The other four 295 onsets were not followed by such prolonged polar-cap flow channel activity, and the

resulting substorm expansions after these onsets were far more limited in duration and inpoleward expansion.

298 The strong poleward expansion and auroral activity that occur during the 299 expansion phase of some substorm, such as the 4 of 8 mentioned above, looks very 300 similar to that following P<sub>dvn</sub> enhancement impacts. This suggests that strong meso-scale 301 flows (common to both) may be a property of the driver of activity for both situations, 302 suggesting that the post P<sub>dvn</sub> enhancement impact activity and extended substorm 303 expansion phase activity may be quite similar phenomena. Another variant of this 304 possibility is shown in Figure 9, and appears quite dramatically in the supplemental 305 movie. Figure 9 shows mergers of images from the THEMIS ASIs overlaid with line-of-306 sight (LOS) flow velocities from the SuperDARN radars for selected times throughout an 307 event that started as a typical substorm onset at ~0803 UT during a storm on 13 October 308 2010. Onset was identified from a brightening and activity along the equatorward-most 309 arc at  $\Lambda \sim 61.5^{\circ}$ . The movie shows the same event at the highest possible time resolution 310 (3 s) of the ASIs (though the time resolution of the radar observations is 1 min). This 311 event was chosen because of the excellent coverage from the Inuvik radar of echoes 312 within the polar cap and within the longitude range of the observed auroral activity.

313 After onset, the expansion phase activity expanded poleward and azimuthally 314 until, at ~0813 UT, it reached the pre-existing arc that lay along the poleward boundary 315 of the auroral oval (at  $\Lambda \sim 66^{\circ}$ , near 21 MLT, and marked by an orange dashed line). The 316 expansion phase aurora subsequently did not protrude further into the pre-existing polar 317 cap region and visible streamers stopped penetrating into the equatorial portion of the 318 oval. Then, at ~0829 UT, the activity enhanced. It began along the auroral poleward 319 boundary, and was followed by poleward expansion to  $\Lambda \sim 70^{\circ}$  and numerous streamers 320 extending to within the equatorward portion of the oval. The radar observations show 321 moderate LOS flows (a few hundred m/s) as the post-onset auroral activity expanded 322 poleward to the pre-existing auroral poleward boundary and became stalled at that 323 boundary. Then, starting at ~0827- 0828 UT, the LOS flow velocities strongly increased 324 to near 1 km/s, and the further auroral activity and poleward expansion commenced at 325 ~0829 UT. Both the flows and activity reduced considerably starting at ~0840 UT. The 326 renewed activity after 0828 UT for this event resembles that from the P<sub>dvn</sub> enhancements 327 in that the activity started from the polar cap boundary, initiated as the flows directed 328 toward the auroral poleward boundary increased, and decreased as the incoming flows 329 decreased. A difference is that the re-enhancement of auroral activity for this event 330 appears to have occurred  $\sim 1-2$  min after the first detection of the polar-cap flow increases, 331 indicating that the enhanced flows may have propagated from the dayside. We suggest 332 that the flow enhancement driven by the shock impact likely occurred approximately 333 simultaneously over the entire polar cap, including adjacent to the auroral poleward 334 boundary. Also, the longitudinal extent of the activity was much smaller for this event 335 than for the P<sub>dvn</sub> enhancement impacts. We suspect that this reflects a smaller longitude 336 range of the enhanced polar cap flows; however, there were not sufficient radar echoes 337 away from the activity region to determine whether or not this was the case.

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## 4. Summary and Conclusions

339 We have used a combination of ground-based ASI, radar, and low-altitude 340 spacecraft observations to further describe the response of the magnetosphere-ionosphere 341 system to abrupt enhancements of  $P_{dyn}$  under southward IMF conditions, such as can 342 occur when CME shocks impact the magnetosphere and lead to magnetic storms. 343 Consistent with previous studies, we find that P<sub>dyn</sub> enhancements promptly lead to 344 enhanced nightside auroral activity and large auroral poleward expansion, which 345 indicates strong nightside reconnection. The auroral activity and poleward expansion 346 appears to be very similar to what occurs during the expansion phase of substorms, and 347 includes numerous auroral streamers, which indicate plasma sheet flow bursts (as directly 348 seen for the 5 April 2010 event). However, the activity is much broader in MLT, has limited azimuthal expansion, and does not show the classical brightening of theequatorward-most arc that demarcates a substorm onset.

351 The radar observations within the nightside polar cap have given possible insight 352 into what is driving this shock-driven auroral response, including the associated 353 enhancement in auroral-oval ionospheric and field-aligned currents and in reconnection 354 activity along the auroral poleward boundary. Specifically, we have found a quick and 355 large enhancement of meso-scale polar cap flows heading towards the nightside 356 separatrix after shock impact. Recent studies have shown that such enhanced meso-scale 357 flows often lead to localized, enhanced flows across the nightside separatrix, (i.e, 358 localized reconnection), and that these flows lead to plasma sheet flow bursts, PBIs, and 359 streamers. Furthermore, PBI observations have shown that the polar cap boundary can 360 move poleward associated with the reconnection that results from the impact onto the 361 plasma sheet of enhanced flows from along open, polar cap field lines.

362 Based on these previous observations, we suggest that the flow enhancements 363 along polar-cap field line that are promptly driven by Pdyn enhancement impacts are an 364 attractive candidate for driving the almost immediate strong auroral, current, and 365 reconnection activity resulting from the impact. These flow enhancement after P<sub>dvn</sub> 366 impact would have to extend to just poleward of the auroral poleward boundary to 367 account for the rapid auroral response of impact. However we are not able to definitively 368 show that this occurred due to the lack of sufficient flow observations adjacent to the 369 aurora boundary at impact. To account for the poleward expansion of the oval that 370 extends over several hours of MLT and brings the auroral poleward boundary poleward 371 by up to several degrees in latitude over a period of several minutes, enhanced meso-372 scale flows would have to impact the auroral poleward boundary over the entire 373 longitudinal extend of the auroral poleward expansion and such impacts must persist 374 throughout the period of the auroral poleward expansion. The observations presented 375 here are consistent with this scenario, the enhanced flows and their persistence during the

376 period of auroral activity and expansion being seen within regions of polar cap radar 377 echoes. However, we do not have radar echoes over the entire longitude extent of the 378 auroral activity. Additionally, the enhancement in field-aligned currents by the shock 379 impact may be an additional contributor to the shock-associated auroral brightening and 380 current enhancement.

381 We have also noted that enhanced meso-scale flows along polar-cap field line 382 occur under other conditions, and they may drive enhanced auroral activity and poleward 383 expansion during these conditions. In particular, we have previously found evidence that 384 prolonged periods of auroral activity and auroral poleward expansion during the 385 expansion phase of substorms accompany prolonged periods of strong polar-cap flow 386 channels directed toward the nightside auroral oval boundary. More limited expansion 387 phase activity was seen when such flow channels where absent following substorm onset. 388 In this paper, we presented an example showing variation of this phenomenon. In the 389 example, auroral activity after a substorm onset continued only for the several minutes 390 that it took for the activity to expand poleward to the pre-existing auroral poleward 391 boundary. Then, after a delay of  $\sim 15$  min, polar cap flows towards the nightside auroral 392 poleward boundary strongly increased to near 1 km/s, and further auroral activity started 393 from along the auroral poleward boundary. This was followed by further auroral 394 poleward expansion and numerous streamers extending to within the equatorward portion 395 of the oval. Both the flows and activity reduced considerably starting at the same time 396 after ~15 min of activity.

397 It is potentially interesting that we have seen common features of, and possibly 398 common driving of, the activity after  $P_{dyn}$  enhancement impacts and following substorm 399 onsets. The common features include auroral activity protruding into the pre-existing 400 polar cap and numerous streamers penetrating towards the equatorial portion of the 401 auroral oval. These common features make activity following  $P_{dyn}$  enhancements under 402 southward IMF look very similar to that in post-substorm onset examples. Furthermore, 403 the activity is associated with enhanced meso-scale flows heading towards the nightside 404 separatrix in both sets of cases, and the enhanced flows and activity ceases essentially 405 simultaneously in both sets. It is also interesting that for the Pdyn enhancement impact 406 events, and the post-substorm event presented here, activity starts from the auroral polar 407 cap boundary and not from the brightening and breakup of an equatorward auroral arc as 408 in the sequence that initiates a substorm. We thus suggest that enhanced meso-scale 409 flows heading toward the nightside polar cap boundary may be common to the driving of 410 auroral activity that extents into the pre-existing polar cap, and that such events may start 411 from the auroral poleward boundary and may not necessarily be associated with a typical 412 substorm onset.

413 If further investigation indicates that above suggestion may be true, it would open 414 up some new potentially significant questions. For example, enhancements in polar cap 415 convection are often viewed as leading to substorm growth phase conditions. Under what 416 conditions do such enhancements lead to a substorm growth phase, versus leading 417 directly to strong poleward expansion of, and strong activity within, the region of auroral 418 oval field lines? Another question is what leads to and causes the enhancements in meso-419 scale polar cap flows. Clearly, enhancements in P<sub>dyn</sub> do so, but how do they do this? 420 Also, what drives the polar-cap flow enhancements when there is not an enhancement in 421 P<sub>dyn</sub>? Southward turnings of the IMF are one obvious possibility, but such a turning was 422 not observed for the example presented here.

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#### 448 **References**

- 449 Baker, D.N., Jaynes, A.N., Li, X., Henderson, M.G., Kanekal, S.G., Reeves, G.D., 450 Spence, H.E., Claudepierre, S.G., Fennell, J.F., Hudson, M.K., Thorne, R.M., Foster, J.C., Erickson, P.J., Malaspina, D.M., Wygant, J.R., Boyd, A., Kletzing, 451 452 C.A., Drozdov, A., Shprits, Y.Y., 2014. Gradual diffusion and punctuated phase 453 space density enhancements of highly relativistic electrons: Van Allen Probes 454 observations. Res. Lett. 41. 2013GL058942. Geophys. doi:10.1002/2013GL058942 455
- Boudouridis, A., 2003. Effect of solar wind pressure pulses on the size and strength of the
   auroral oval. J. Geophys. Res. 108. doi:10.1029/2002JA009373
- Chua, D., Parks, G., Brittnacher, M., Peria, W., Germany, G., Spann, J., Carlson, C.,
  2001. Energy characteristics of auroral electron precipitation: A comparison of
  substorms and pressure pulse related auroral activity. J. Geophys. Res. Space
  Phys. 106, 5945–5956. doi:10.1029/2000JA003027
- de la Beaujardière, O., Lyons, L.R., Ruohoniemi, J.M., Friis-Christensen, E., Danielsen,
  C., Rich, F.J., Newell, P.T., 1994. Quiet-Time Intensifications Along the
  Poleward Auroral Boundary Near Midnight. J. Geophys. Res. 99, 287–298.
  doi:10.1029/93JA01947
- 466 Durgonics, T., Komjathy, A., Verkhoglyadova, O., Shume, E.B., Benzon, H.-H.,
  467 Mannucci, A.J., Butala, M.D., Høeg, P., Langley, R.B., 2017. Multiinstrument
  468 observations of a geomagnetic storm and its effects on the Arctic ionosphere: A
  469 case study of the 19 February 2014 storm. Radio Sci. 52, 2016RS006106.
  470 doi:10.1002/2016RS006106
- 471 Gjerloev, J.W., 2012. The SuperMAG data processing technique. J. Geophys. Res. Space
   472 Phys. 117, A09213. doi:10.1029/2012JA017683
- Gkioulidou, M., Ukhorskiy, A., Mitchell, D.G., Sotirelis, T., Mauk, B., Lanzerotti, L.J.,
  2014. The role of small-scale ion injections in the buildup of Earth's ring current
  pressure: Van Allen Probes observations of the March 17th, 2013 storm. J.
  Geophys. Res. Space Phys. 2014JA020096. doi:10.1002/2014JA020096
- Hudson, M.K., Paral, J., Kress, B.T., Wiltberger, M., Baker, D.N., Foster, J.C., Turner,
  D.L., Wygant, J.R., 2015. Modeling CME-shock-driven storms in 2012–2013:
  MHD test particle simulations. J. Geophys. Res. Space Phys. 120, 2014JA020833.
  doi:10.1002/2014JA020833
- Juusola, L., Østgaard, N., Tanskanen, E., Partamies, N., Snekvik, K., 2011. Earthward
  plasma sheet flows during substorm phases. J. Geophys. Res. Space Phys. 116,
  A10228. doi:10.1029/2011JA016852
- Kokubun, S., McPherron, R.L., Russell, C.T., 1977. Triggering of substorms by solar
   wind discontinuities. J. Geophys. Res. 82, 74–86. doi:10.1029/JA082i001p00074
- Koustov, A.V., St.-Maurice, J.-P., Sofko, G.J., Andre, D., MacDougall, J.W., Hairston,
  M.R., Fiori, R.A., Kadochnikov, E.E., 2009. Three-way validation of the Rankin
  Inlet PolarDARN radar velocity measurements. Radio Sci. 44.
  doi:10.1029/2008RS004045

- Lee, D.-Y., Lyons, L.R., 2004. Geosynchronous magnetic field response to solar wind dynamic pressure pulse. J. Geophys. Res. Space Phys. 109, A04201. doi:10.1029/2003JA010076
- Lee, D.-Y., Lyons, L.R., Reeves, G.D., 2005. Comparison of geosynchronous energetic
  particle flux responses to solar wind dynamic pressure enhancements and
  substorms. J. Geophys. Res. Space Phys. 110, A09213.
  doi:10.1029/2005JA011091
- Li, W., Thorne, R.M., Ma, Q., Ni, B., Bortnik, J., Baker, D.N., Spence, H.E., Reeves,
  G.D., Kanekal, S.G., Green, J.C., Kletzing, C.A., Kurth, W.S., Hospodarsky, G.B.,
  Blake, J.B., Fennell, J.F., Claudepierre, S.G., 2014. Radiation belt electron
  acceleration by chorus waves during the 17 March 2013 storm. J. Geophys. Res.
  Space Phys. 119, 2014JA019945. doi:10.1002/2014JA019945
- Li, Z., Hudson, M., Kress, B., Paral, J., 2015. Three-dimensional test particle simulation
  of the 17–18 March 2013 CME shock-driven storm. Geophys. Res. Lett. 42,
  2015GL064627. doi:10.1002/2015GL064627
- Liou, K., Newell, P.T., Meng, C.-I., Wu, C.-C., Lepping, R.P., 2003. Investigation of
  external triggering of substorms with Polar ultraviolet imager observations. J.
  Geophys. Res. Space Phys. 108, 1364. doi:10.1029/2003JA009984
- Loto'aniu, T.M., Singer, H.J., Rodriguez, J.V., Green, J., Denig, W., Biesecker, D.,
  Angelopoulos, V., 2015. Space weather conditions during the Galaxy 15
  spacecraft anomaly. Space Weather 13, 2015SW001239.
  doi:10.1002/2015SW001239
- Lyons, L.R., 2005. Global auroral responses to abrupt solar wind changes: Dynamic
   pressure, substorm, and null events. J. Geophys. Res. 110.
   doi:10.1029/2005JA011089
- Lyons, L.R., Gallardo-Lacourt, B., Zou, S., Weygand, J.M., Nishimura, Y., Li, W.,
  Gkioulidou, M., Angelopoulos, V., Donovan, E.F., Ruohoniemi, J.M., Anderson,
  B.J., Shepherd, S.G., Nishitani, N., 2016. The 17 March 2013 storm: Synergy of
  observations related to electric field modes and their ionospheric and
  magnetospheric Effects. J. Geophys. Res. Space Phys. 121, 2016JA023237.
  doi:10.1002/2016JA023237
- Lyons, L.R., Nishimura, Y., Kim, H.-J., Donovan, E., Angelopoulos, V., Sofko, G.,
  Nicolls, M., Heinselman, C., Ruohoniemi, J.M., Nishitani, N., 2011. Possible
  connection of polar cap flows to pre- and post-substorm onset PBIs and streamers.
  J. Geophys. Res. 116, 14 PP. doi:201110.1029/2011JA016850
- Lyons, L.R., Zesta, E., Samson, J.C., Reeves, G.D., 2000. Auroral disturbances during
  the January 10, 1997 magnetic storm. Geophys. Res. Lett. 27, PP. 3237-3240.
  doi:200010.1029/1999GL000014
- McPherron, R.L., 1970. Growth phase of magnetospheric substorms. J. Geophys. Res. 75,
   5592–5599. doi:10.1029/JA075i028p05592
- Mende, S.B., Harris, S.E., Frey, H.U., Angelopoulos, V., Russell, C.T., Donovan, E.,
   Jackel, B., Greffen, M., Peticolas, L.M., 2008. The THEMIS Array of Ground-

- based Observatories for the Study of Auroral Substorms. Space Sci. Rev. 141,
  357–387. doi:10.1007/s11214-008-9380-x
- Nishimura, Y., Lyons, L.R., Zou, S., Xing, X., Angelopoulos, V., Mende, S.B., Bonnell,
  J.W., Larson, D., Auster, U., Hori, T., Nishitani, N., Hosokawa, K., Sofko, G.,
  Nicolls, M., Heinselman, C., 2010. Preonset time sequence of auroral substorms:
  Coordinated observations by all-sky imagers, satellites, and radars. J. Geophys.
  Res. 115. doi:10.1029/2010JA015832
- Ohtani, S., Yoshikawa, A., 2016. The initiation of the poleward boundary intensification
  of auroral emission by fast polar cap flows: A new interpretation based on
  ionospheric polarization. J. Geophys. Res. Space Phys. 121, 2016JA023143.
  doi:10.1002/2016JA023143
- Paxton, L., Meng, C., Fountain, G., Ogorzalek, B., Darlington, E., Gary, S., Goldsten, J.,
  Kusnierkiewicz, D., Lee, S., Linstrom, L., Maynard, J., Peacock, K., Persons, D.,
  Smith, B., 1992. Special Sensor Ultraviolet Spectrographic Imager (ssusi) an
  Instrument Description. Spie Int Soc Optical Engineering, Bellingham.
- Paxton, L.J., Christensen, A.B., Humm, D.C., Ogorzalek, B.S., Pardoe, C.T., Morrison,
  D., Weiss, M.B., Crain, W., Lew, P.H., Mabry, D.J., Goldsten, J.O., Gary, S.A.,
  Persons, D.F., Harold, M.J., Alvarez, E.B., Ercol, C.J., Strickland, D.J., Meng, C.I., 1999. Global ultraviolet imager (GUVI): measuring composition and energy
  inputs for the NASA Thermosphere Ionosphere Mesosphere Energetics and
  Dynamics (TIMED) mission. pp. 265–276. doi:10.1117/12.366380
- Paxton, L.J., Schaefer, R.K., Zhang, Y., Kil, H., 2017. Far ultraviolet instrument technology. J. Geophys. Res. Space Phys. 122, 2016JA023578. doi:10.1002/2016JA023578
- Paxton, L.J., Spisz, T., Crowley, G., Gary, R., Hopkins, M.M., Morrison, D., Wiess, M.,
  Fountain, G.H., Suther, L., Meng, C.-I., Strickland, D.J., 1998. Interactive
  interpretation and display of far Ultraviolet Data. Adv. Space Res. 22, 1577–1582.
  doi:10.1016/S0273-1177(99)00116-7
- Pitkänen, T., Aikio, A.T., Juusola, L., 2013. Observations of polar cap flow channel and
   plasma sheet flow bursts during substorm expansion. J. Geophys. Res. Space Phys.
   118, 774–784. doi:10.1002/jgra.50119
- Shi, Y., Zesta, E., Lyons, L.R., Yang, J., Boudouridis, A., Ge, Y.S., Ruohoniemi, J.M.,
  Mende, S., 2012. Two-dimensional ionospheric flow pattern associated with
  auroral streamers. J. Geophys. Res. 117. doi:10.1029/2011JA017110
- Waters, C.L., Anderson, B.J., Liou, K., 2001. Estimation of global field aligned currents
   using the iridium® System magnetometer data. Geophys. Res. Lett. 28, 2165–
   2168. doi:10.1029/2000GL012725
- Yue, C., Nishimura, Y., Lyons, L.R., Angelopoulos, V., Donovan, E.F., Shi, Q., Yao, Z.,
  Bonnell, J.W., 2013. Coordinated THEMIS spacecraft and all-sky imager
  observations of interplanetary shock effects on plasma sheet flow bursts,
  poleward boundary intensifications, and streamers. J. Geophys. Res. Space Phys.
  118, 3346–3356. doi:10.1002/jgra.50372

- Zesta, E., Singer, H.J., Lummerzheim, D., Russell, C.T., Lyons, L.R., Brittnacher, M.J.,
  2000. The Effect of the January 10, 1997, Pressure Pulse on the MagnetosphereIonosphere Current System, in: Ohtani, S.-I., Fujii, R., Hesse, M., Lysak, R.L.
  (Eds.), Magnetospheric Current Systems. American Geophysical Union, pp. 217–
  226.
- Zhou, X., Tsurutani, B.T., 2001. Interplanetary shock triggering of nightside geomagnetic
  activity: Substorms, pseudobreakups, and quiescent events. J. Geophys. Res.
  Space Phys. 106, 18957–18967. doi:10.1029/2000JA003028
- Zou, Y., Nishimura, Y., Lyons, L.R., Donovan, E.F., Ruohoniemi, J.M., Nishitani, N.,
  McWilliams, K.A., 2014. Statistical relationships between enhanced polar cap
  flows and PBIs. J. Geophys. Res. Space Phys. 119, 2013JA019269.
  doi:10.1002/2013JA019269
- Zou, Y., Nishimura, Y., Lyons, L.R., Donovan, E.F., Shiokawa, K., Ruohoniemi, J.M., McWilliams, K.A., Nishitani, N., 2015. Polar Cap Precursor of Nightside Auroral Oval Intensifications Using Polar Cap Arcs. J. Geophys. Res. Space Phys. 2015JA021816. doi:10.1002/2015JA021816

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- Figure 1: From top to bottom, the WIND  $P_{dyn}$ , the OMNI IMF, the SuperMAG upper U and lower L magnetic index, and the SuperMAG ring current index for all MLT and within the dusk, noon, dawn, and midnight sectors on 17 March 2013. At the bottom are magnetic perturbations observed along Iridium satellite trajectories during the two 10 min intervals identified by the purple arrows. Red and blue shadings give upward and downward radial current, respectively, obtained from the curl of fits to the magnetic perturbations (based on Figurers 1 and 6 of Paper 1).
- Figure 2: Representative mergers of auroral images over Canada from the array of THEMIS ASIs for the shock impact event on 17 March 2013 overlaid with LOS flow velocities from the SuperDARN radars. UT times are given in the upper left comer of each panel. LOS flow velocities within the polar cap are primarily from Inuvik to the west and Rankin Inlet to the east. There are also some LOS velocities from the from lower latitude SuperDARN radars. Yellow ellipses encircle the enhanced anti-sunward LOS flows seen within the polar cap by the radars after the shock impact.
- Figure 3: Auroral images over the southern hemisphere from the Special Sensor Ultraviolet Spectrographic Imager onboard polar-orbiting DMSP spacecraft at the LBH2 (165-180 nm) (135.6 nm/) band. The images are essentially keograms made along space trajectories from horizon to horizon imager scans in the direction normal to the trajectory. Precipitating electron energy fluxes measured along the DMSP trajectories are color coded along the trajectories. Flows measures by the DMSP F17 and F16 spacecraft in the direction normal to the spacecraft trajectory are shown by violet bars normal along the trajectories. UT times indicated along each trajectory.
- Figure 4: (modify to make SuperDARN flows scales the same) Precipitating electron energy fluxes from both the POES (left panel) and DMSP (right panel) spacecraft

over the southern hemisphere. Polar cap flows seen by the southern hemisphere SuperDARN radars at times during the period of rapid auroral poleward expansion are also shown, as are the flows measures by the DMSP F17 and F16 spacecraft in the direction normal to the spacecraft trajectory (black bars along the trajectories). UT times indicated along each trajectory

- Figure 5: From top to bottom, the OMNI P<sub>dyn</sub>, the OMNI IMF, the SuperMAG upper U and lower L magnetic index, and the SuperMAG ring current index for all MLT and within the dusk, noon, dawn, and midnight sectors on 19 March 2014.
- Figure 6: Same as Figure 2, except for 19 February 2014.
- Figure 7: SuperDARN LOS flows in the southern hemisphere for times before and after the shock impact on 19 February 2014. Precipitating electron energy fluxes from the POES spacecraft that are shown in the fourth panel, with UT times indicated along each trajectory.
- Figure 8: SuperDARN LOS flows in the northern hemisphere for times before after the shock impact on 5 April 2015.
- Figure 9: Mergers of images from the THEMIS ASIs overlaid with LOS flow velocities from the SuperDARN radars for selected times during an event that started as a typical substorm onset at ~0803 UT on 13 October 2010.
- Supplemental Movie: Mergers of images from the THEMIS ASIs overlaid with LOS flow velocities from the SuperDARN radars at the high possible time resolution (3 s) of the ASIs during the event that started as a typical substorm onset at ~0803 UT on 13 October 2010. Time resolution of the radar observations is 2 min.

















